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Supplemental Online Material for

## Ocean and Coastal Acidification off New England and Nova Scotia

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## Supplemental Methods

### Primary Controls on NECAN Seasonal Dynamics

Time-series analyses were evaluated at each of the locations denoted in Figure 4 using actual discrete time-series data where available (western Gulf of Maine, Long Island Sound) or by extracting estimates using the gridded data product produced by Signorini et al. (2013).

### Signorini-Based Estimates

Figure 4 denotes the estimated individual effects of CO<sub>2</sub> solubility (SOL), air-sea CO<sub>2</sub> flux (AS), mixing (MIX) and biologic activity (BIO) on  $\Omega_a$  at six New England/Nova Scotia locations. . . These estimates were derived from the calculations made using modeled 12-month climatologies of sea surface temperature (SST), sea surface salinity (SSS), total alkalinity (TA),  $p\text{CO}_{2,\text{air-sea}}$  CO<sub>2</sub> flux (FLUX), and surface mixed-layer depth (MLD). BIO was determined by combining December TA,  $p\text{CO}_{2,\text{sea}}$ , SST, and SSS to derive dissolved inorganic carbon ( $\text{DIC}_{\text{TA-}p\text{CO}_2}$ ) using the CO2SYS program (Lewis and Wallace, 1998;  $K_1$  and  $K_2$  of Millero, 2010;  $\text{KSO}_4$  of Dickson, 1990; TB [total boron] of Uppstrom, 1974), then deriving  $p\text{CO}_{2,\text{sea,TA-DIC}}$  and  $\Omega_{\text{arag,TA-DIC}}$  at monthly SST and SSS using December TA and  $\text{DIC}_{\text{TA-}p\text{CO}_2}$ . AS was determined by adding (for release of CO<sub>2</sub> to the atmosphere) or removing (for uptake of CO<sub>2</sub> from the atmosphere) the DIC (dissolved inorganic carbon) represented by each monthly FLUX out of or into the MLD ( $\text{FLUX}_{\text{DIC}}$ ) to the DIC calculated for the previous month, then deriving  $p\text{CO}_{2,\text{sea}}$  and  $\Omega_{\text{arag}}$  from monthly TA and  $\text{DIC} + \text{FLUX}_{\text{DIC}}$ . AS MLD depth was held to at least 5 m, which is was arbitrarily chosen as a logical minimum depth over which CO<sub>2</sub> is added or removed. We note that during the strongly stratified summer season, 5 m can be 1–3 m less than the pycnocline. MIX was determined using the Gulf of Maine TA-SSS and DIC-SSS using regressions taken

from Table 1 of Wang et al. (2013) using a TA:DIC ratio of 1:07  $TA = 37.3*SSS + 998$ ;  $DIC = 34.6*SSS + 933$ . Monthly TA and DIC were calculated according to these regressions ( $TA_{Wang}$  and  $DIC_{Wang}$ , respectively), with changes from month n-1 to month n calculated as  $\Delta TA_{Wang-n} = TA_{Wang-n} - TA_{Wang-n-1}$  and  $\Delta DIC_{Wang-n} = DIC_{Wang-n} - DIC_{Wang-n-1}$ . Then MIX  $pCO_2$  and  $\Omega_a$  were derived from monthly SSS together with TA and DIC at month n:  $TA_n = TA_{n=0} + \Delta TA_{Wang-n}$  and  $DIC_n = DIC_{n=0} + \Delta DIC_{Wang-n}$ . BIO was calculated as the residual between the climatologies of  $pCO_2$  and  $\Omega_a$  and the SOL, AS and MIX terms:  $BIO_{pCO_2} = pCO_2 - [(SOL - pCO_2) + (AS - pCO_2) + (MIX - pCO_2)]$ .

### Western Gulf of Maine Data

The calculated individual effects on  $pCO_{2,sw}$  and  $\Omega_a$  of changes in  $CO_2$  solubility (SOL), air-sea  $CO_2$  flux (AS), and mixing (MIX) at UNH Buoy D were calculated as above. SSS, SST,  $pCO_2$ , and FLUX source data were monthly climatologies derived from 2006–2014 buoy observations. MLD source data was a monthly climatology derived from 2004–2014 shipboard salinity and temperature profiles ( $n = 151$ ) in the region of UNH Buoy D, again limited to 5 m or deeper. TA was derived from salinity according to a locally derived regression:  $TA = (SSS*52.24) + 476.3$  (unpublished data from author Joe Salisbury).

### Long Island Sound Data

The calculated individual effects on  $pCO_2$  and  $\Omega_{arag}$  of changes in  $CO_2$  solubility (SOL), air-sea  $CO_2$  flux (AS), and mixing (MIX) in Long Island Sound (LIS) were calculated in the *Signorini-based estimate*. SSS, SST, and TA source data were monthly climatologies. A monthly climatology of pH (NBS scale), paired with TA, SSS, and SST was used with CO2SYS as above to generate a monthly  $pCO_2$  climatology. FLUX was calculated from  $pCO_2$ , atmospheric  $pCO_2$

at Mauna Loa (398  $\mu\text{atm}$ ), and monthly winds, according to the k660 parameterization of Ho et al. (2006). MLD for LIS was from the same model employed in the *Signorini-based estimate* at location 41.25°N –71.25°W, and again limited to 5 m or deeper.

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