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# CO<sub>2</sub> Transport in Deep Waters Off Wilkes Land

BY TARO TAKAHASHI AND DAVID W. CHIPMAN

The densest waters in the world ocean are formed on the continental shelf areas around Antarctica by ice formation and by the loss of heat and water to the air. The Weddell Sea (e.g., Gordon, 1971; Carmack and Foster, 1975; Foldvik et al., 1985), the Ross Sea (e.g., Jacobs et al., 1970, 1985) and other continental margin areas such as the coasts of the Adélie and Wilkes Lands (e.g., Gordon and Tchernia, 1972; Carmack and Killworth, 1978; Foster, 1995) are known to be the major producers of these dense waters. These waters spill over the shelf edge and, during their descent, mix with components of Circumpolar Deep Water (CPDW) to form the Antarctic Bottom Water (AABW) that fills the major abyssal basins of the world ocean. This process constitutes an important pathway between the atmosphere and the abyssal ocean for long-term storage of CO<sub>2</sub>. However, the CO<sub>2</sub> transport processes associated with bottom-water formation have not been well documented. Here, we present observations made during the 1992 World Ocean Circulation Experiment (WOCE) S4P cruise showing that atmospheric CO<sub>2</sub> taken up by Wilkes Land shelf waters is transported into the upper AABW.

Figure 1a–d shows four properties along a section (68°S, 173°E–70.7°S, 168°E) that is nearly at right angles to the continental shelf off the coast of Wilkes Land west of the Ross Sea. Waters with the partial pressure of CO<sub>2</sub> (*p*CO<sub>2</sub>) less than the 1992 atmospheric value of 375 μatm are found from the surface to about 200 m depth near the shelf edge (Figure 1b). These waters are likely a mixture of the offshore surface layer and shelf waters that are modified by processes involving sea ice and glacial ice (Jacobs et al., 1985); their low *p*CO<sub>2</sub> values are due to summertime phytoplankton blooms on the shelf and offshore. These waters absorb CO<sub>2</sub> from the atmosphere because of their low *p*CO<sub>2</sub> values (below the atmospheric value).

Circumpolar Deep Water (CPDW), which is characterized by high values of temperature

(> 1.0°C), *p*CO<sub>2</sub>, and total CO<sub>2</sub> concentration (TCO<sub>2</sub>), and by lower silica concentrations, is located immediately under the 100 m thick offshore mixed layer (Figure 1a–d). It appears to mix with the low *p*CO<sub>2</sub> surface water, forming a transition in the shelf break zone down to about 500 m, as described by Gordon (1971). The AABW, with higher concentrations of silica and TCO<sub>2</sub>, is found below the CPDW (Figure 1c,d). A layer of sub-0°C water with lower silica (< 100 μmol kg<sup>-1</sup>) is observed from the surface to 3,000 m along the shelf slope. Near the base of the slope (3,000–3,500 m) lies a pool of dense cold water, the Ross Sea Bottom Water, with lower values of silica, TCO<sub>2</sub>, and *p*CO<sub>2</sub> (Gordon and Tchernia, 1972; Carmack and Killworth, 1978), flowing westward along the base of the continental slope.

The potential temperature-salinity plot in Figure 1e shows the mixing relationships among these water masses. Between the less-cold CPDW and the colder high-salinity Ross Sea Bottom Water, a lower-salinity kink in the linear trend is seen at about 0.2°C and a salinity of 34.67 (sigma-2 density of 37.14 kg m<sup>-3</sup>, about 1,060 m depth). This water has a local minimum for Apparent Oxygen Utilization (AOU) (not shown) and *p*CO<sub>2</sub> values (Figure 1f), indicating a more recent origin. These values, however, are considerably greater than the surface water values due to mixing with the high-AOU, high-*p*CO<sub>2</sub> CPDW. Because of the similarity in the temperature-salinity relationships and in the density, we consider that this water is the same type as that reported by Carmack and Killworth (1978) between 147°E and 162°E: potential temperature of 0° to -0.5°C, salinity of 34.68–34.70 PSS, and sigma-2 density of 37.20–37.25 kg m<sup>-3</sup> at depths of 1,000–2,200 m. Using a turbulent plume model, they demonstrated that this water was formed by sinking of Wilkes Land shelf water that had entrained CPDW during its descent along the continental slope. The water was not dense enough to reach the bottom, but it interleaved into the

thick body of the AABW. Because the shelf water had absorbed atmospheric CO<sub>2</sub> during the summer, the process described above suggests a transport pathway for atmospheric CO<sub>2</sub> into the AABW regime. The Ross Sea Bottom Water located at the base of the continental slope, which had absorbed atmospheric CO<sub>2</sub> when exposed at the surface during summer low *p*CO<sub>2</sub> periods (e.g., Bates et al., 1998; Hales and Takahashi, 2004), transported atmospheric CO<sub>2</sub> into the deepest regime of the AABW.

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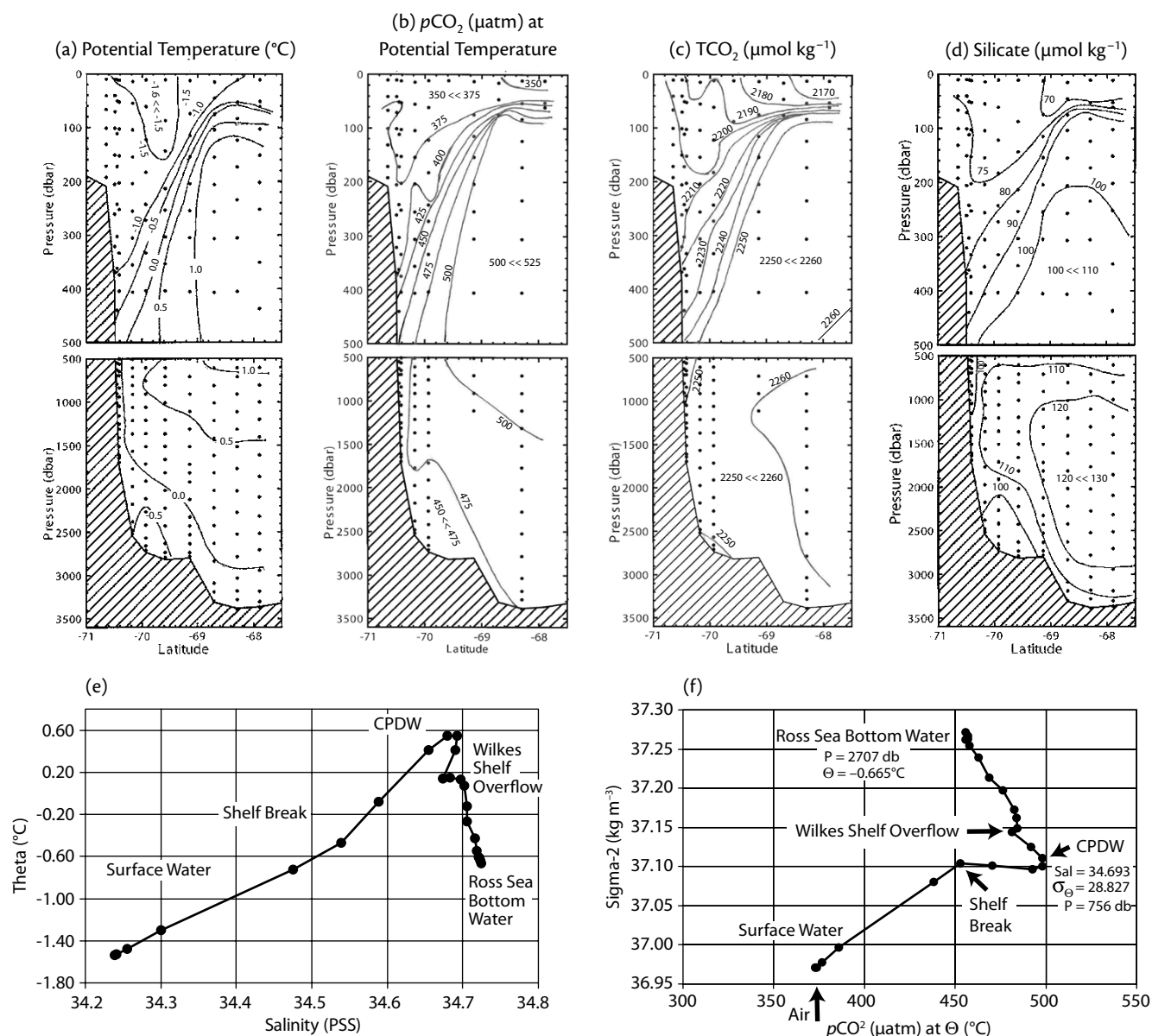


Figure 1. Depth profiles of (a) potential temperature ( $^{\circ}\text{C}$ ), (b)  $p\text{CO}_2$  ( $\mu\text{atm}$  at potential temperature), (c)  $\text{TCO}_2$  ( $\mu\text{mol kg}^{-1}$ ), and (d) silicate concentration ( $\mu\text{mol kg}^{-1}$ ) over the continental slope off the coast of Wilkes Land ( $67^{\circ}03'\text{S}$ ,  $174^{\circ}19'\text{E}$  to  $70^{\circ}39'\text{S}$ ,  $168^{\circ}00'\text{E}$ ) were measured from March 21 to 23, 1992, during the austral summer World Ocean Circulation Experiment (WOCE) S4P cruise (Chipman et al., 1997). Property-property plots are for the data at WOCE S4P Station 775 ( $69^{\circ}56'\text{S}$  and  $169^{\circ}20'\text{E}$ ): (e) potential temperature vs. salinity, and (f) sigma-2 density vs.  $p\text{CO}_2$  at potential temperature. The presence of Wilkes Land shelf water is indicated as Wilkes Land Overflow (OF).

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